

1 **Probabilistic Seismic Hazard Assessment of Itanagar, Arunachal Pradesh,**
2 **India: Insights into Tectonic Activity, Risk Zones, and Hazard Mapping**

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Probabilistic Seismic Hazard Assessment of Itanagar, Arunachal Pradesh, India: Insights into Tectonic Activity, Risk Zones, and Hazard Mapping

Abstract

The paper presents a probabilistic assessment of seismic hazards for the Itanagar region of Arunachal Pradesh, India. In this study, earthquake data is compiled from the United States Geological Survey (USGS) around Itanagar in a circular enclosure of 500 km radius. The catalog is homogenized into a unified scale of moment magnitude. The earthquake data is collected between 1900 and 2024. The Seismotectonic Atlas (SEISAT) provided fault information, which is combined with earthquake information to facilitate detailed analysis and visualization using ArcGIS software. There are 33 active tectonic features in the study area, of which 18 are found to be potential sources of seismic activity. The Gutenberg–Richter (G–R) relationship is used to determine the seismicity parameters for each source zone. The region is divided into four primary subzones based on seismic activity and tectonic characteristics. Based on linear sources identified within and around the study area, this study estimates the seismic hazard of the region. Based on the Probabilistic Seismic Hazard Analysis (PSHA) method, peak ground acceleration (PGA) values indicate a 2% probability of exceeding 0.22g and a 10% probability of exceeding 0.36g over 50 years. A spectral acceleration (S_a) is also assessed for return intervals of 475 years and 2475 years, across 0.1 s, 0.3 s, 0.5 s, 1.0 s, 2.0 s, and 3.0 s. The findings from the study are compared with the with other localities in the Northeast region, as well as with specifications outlined in the IS 1893-Part-1 (2016). The results of this study can be used to develop risk reduction strategies, risk acceptance criteria, and financial analyses based on the results of the comprehensive analysis and higher resolution hazard mapping.

Keywords: Seismic hazard analysis; Earthquake catalogue; Completeness analysis; Peak horizontal acceleration; Spectral acceleration.

1. Introduction

Large earthquakes over the years (such as 1964 Nigatta earthquake, 1897 Assam earthquake, 1934 Bihar-Nepal earthquake, 1950 Assam earthquake, 1988 Bihar-Nepal earthquake, 1989 Loma Prieta earthquake, 1995 Japan earthquake, 1999 Chi-Chi earthquake, 2010 Chile earthquake, 2011 Japan

51 earthquake, 2015 Nepal earthquake, and 2023 Turkey earthquake) have left many lessons to ponder
52 and understand in order to develop preventative measures and regulations to reduce the future
53 tragedies. Since the prediction of damages or destructions caused by an earthquake in any specific
54 area depends on several factors such as seismicity and topography of the area, type and condition
55 of subsurface soil, groundwater and intensity of shaking, it is necessary to take the proper steps to
56 assess the seismic hazards in order to obtain the precise estimations of seismic hazard parameters
57 [1-6]. Apart from the aforementioned factors, some other influencing factors of seismic hazards at
58 any particular location also account the magnitude of earthquake, duration of ground shaking,
59 source to site distance, and the return period [7]. Therefore, the estimation of hazard analysis,
60 considering aforementioned parameters, caused by such a large earthquake must be carried out to
61 safe design of critical structures such as high-rise buildings, bridges and highways, which can be
62 done either by adopting deterministic seismic hazard analysis (DSHA) and Probabilistic seismic
63 hazard analysis (PSHA). DSHA estimates the strong ground motion parameters considering worst-
64 case scenario earthquake, i.e., maximum credible earthquake that may severely affect the region,
65 at a distance close to the site of interest. Since earthquake and site of interest contains several
66 uncertainties, which is not accounted in DSHA, the estimated parameters can be best suitable for
67 the project that does not requires great level of accuracy. Therefore, to get more accurate
68 parameters with high level of accuracy considering probability of the occurrence of different
69 earthquakes and the associated uncertainties, PSHA is the best option to estimate the earthquake
70 resistant design parameters accurately [8]. Moreover, PSHA expresses the risk parameters
71 numerically based on the correlation between the characteristics of the local seismic attenuation
72 and the probability of occurrence. As a function of return period and fault displacement, it
73 determines the likelihood that a site will exceed a predetermined ground motion level. Due to its
74 ability to incorporate uncertainty, PSHA is widely used in seismic hazard studies [9].

75

76 Seismic hazard assessments across different regions of India reveal different degrees of seismic
77 risks, as is reported by several researchers [10-24]. A seismic hazard assessment was conducted
78 for the Himalayas and surrounding areas by Shanker and Sharma [25], focusing on the region
79 located between 20° N - 36° N latitude and 69° E - 100° E longitude. From 1900 to 1990,
80 earthquake data were collected from the Himalayan region and were divided into six seismogenic
81 zones, with b values between 0.58 and 1.52. However, the research remained from performing the

82 completeness study of earthquake catalog. Further, a microzonation study for Delhi region was
83 carried out by Iyengar and Ghosh [26], with a 300 km radius centred on India Gate, using PSHA
84 and reported PGA of 0.2g. Anbazhagan *et al.* [27] have conducted PSHA for the Bangalore region
85 considering low-to-moderate seismic hazard zone, and the estimated PGA for this region was
86 found to be 0.121g. This obtained PGA from PSHA is slightly lower than that obtained from
87 DSHA; however, this is higher than the value reported in global seismic hazard maps. To refine
88 the seismic hazard assessment, consideration of seismogenic sources within the radius of 350 km
89 proposed by Gutenberg-Richter and Kijko-Sellevoll emphasized the importance of updating
90 hazard maps and building codes to reflect local risks [27]. Shukla and Choudhury [28] assessed
91 the probabilistic seismic hazard and estimated the site specific ground motions for the Kandla and
92 Mundra ports located in the Gulf of Kachchh to decipher that the site amplification factor varied
93 between 1.37 to 1.94. Further, to design a critical infrastructure like Kakrapar Atomic Power
94 Station in Gujarat, a PSHA study has been performed by Mohanty and Verma [29] and PGA was
95 found to be 0.23g considering maximum credible earthquake with notable contributions from the
96 Narmada-Tapti and Rann of Kutch regions. This study also emphasized both maximum credible
97 and design-basis earthquake scenarios to ensure the resilience of essential facilities. Ashish *et al.*
98 [30] conducted a study to identify areas of high seismic risk in Peninsular India. Despite the region
99 being characterized by a stable continental crust with moderate seismic activity, their PSHA
100 estimated a PGA of 0.4g for a 10% exceedance probability over a 50-year period. Desai and
101 Choudhury [31] identified the spatially varying probabilistic seismic hazard in the Mumbai region.
102 It was deciphered that the codal provisions of IS-1893 [32], which depend on the non-probabilistic
103 seismic hazard assessment, underestimate the potential seismic hazard of the entire Mumbai city
104 especially the Navi Mumbai region that exhibited a significantly high probabilistic seismic hazard.
105 To safeguard cultural heritage structures like the Gol Gumbaz in Vijayapura, South India, Patil *et*
106 *al.* [33] carried out a PSHA study. The study reported PGA of 0.074g and 0.142g for 10% and 2%
107 exceedance probabilities over a 50-year period, respectively. These findings emphasize the
108 necessity of implementing seismic protection measures to preserve such valuable heritage sites.
109 Shukla and Solanki [34] conducted a PSHA study for Indore city, compiling an earthquake catalog
110 within a 400 km radius, which included the 1997 Jabalpur earthquake. The study provided hazard
111 maps with peak ground acceleration (PGA) estimates for 10% and 2% exceedance probabilities
112 over a 50-year period. Thus, it can be stated that region-specific seismic hazards analysis to

113 safeguard infrastructure, cultural sites, heritage structures and communities considering all
114 possible uncertainties associated with earthquakes and the region across India.

115
116 Northeast India is one of the most seismically active regions in the world, experiencing several
117 large earthquakes such as the 1897 Shillong earthquake (M_w 8.1), 1950 Assam earthquake (M_w 8.7),
118 2011 Sikkim earthquake (M_w 6.9), 2015 Nepal earthquake (M_w 7.8), 2016 Manipur Earthquake
119 (M_w 6.7), 2017 Tripura earthquake (M_w 5.7), 2020 Mizoram earthquake (M_w 5.6), 2021 Assam
120 earthquake (M_w 6.4), 2022 Arunachal Pradesh earthquake (M_w 5.7), 2023 Meghalaya (M_w 5.4),
121 2025 Manipur earthquake (M_w 5.7). Therefore, a study has been conducted for Northeast India
122 region using PSHA by Das *et al.* [35], dividing this region into nine seismogenic source zones to
123 capture the local variations in tectonic characteristics. This study estimates PGA and S_a values at
124 bedrock level for return periods of 100, 225, 475, 2475, and 10000 years. The findings revealed
125 that the seismic hazard is underestimated in some areas as per current Indian seismic code. Further,
126 PSHA studies has been done for Tripura and Mizoram region of Northeast India by Sil *et al.* [36],
127 covering a catalog of earthquake events dating back to 1731 within a radius of 500 km from the
128 state boundaries, and provided seismic hazard curves, PGA, and uniform hazard spectra for the
129 region. This study also revealed the spatial variations of peak horizontal acceleration (PHA) with
130 probabilities of exceedance at 2% and 10% in 50 years, emphasizing the importance of updated
131 hazard assessments in this seismically active region.

132
133 Past studies conducted by various researchers [35-37] suggested that the existing seismic hazard
134 maps and Indian seismic codes may underestimate or inadequately capture the seismic risk in
135 specific localized regions. For instance, Das *et al.* [35] indicated that current Indian seismic codes
136 potentially underestimate seismic hazards in Northeastern region of India. Furthermore, previous
137 research by Sil *et al.* [36] and Borgohain *et al.* [38] emphasized the significant spatial variability
138 in ground motion within Northeast India, necessitating detailed local studies to ensure accurate
139 and effective seismic hazard estimations. Consequently, this study was undertaken specifically for
140 Itanagar, Arunachal Pradesh located in close proximity to the Himalayan plate boundary to provide
141 region-specific seismic hazard assessment, enabling effective earthquake-resistant structural
142 design and risk mitigation planning. Therefore, the necessity of conducting a comprehensive
143 PSHA in Itanagar arises to address the variability in ground motion amplification due to local soil

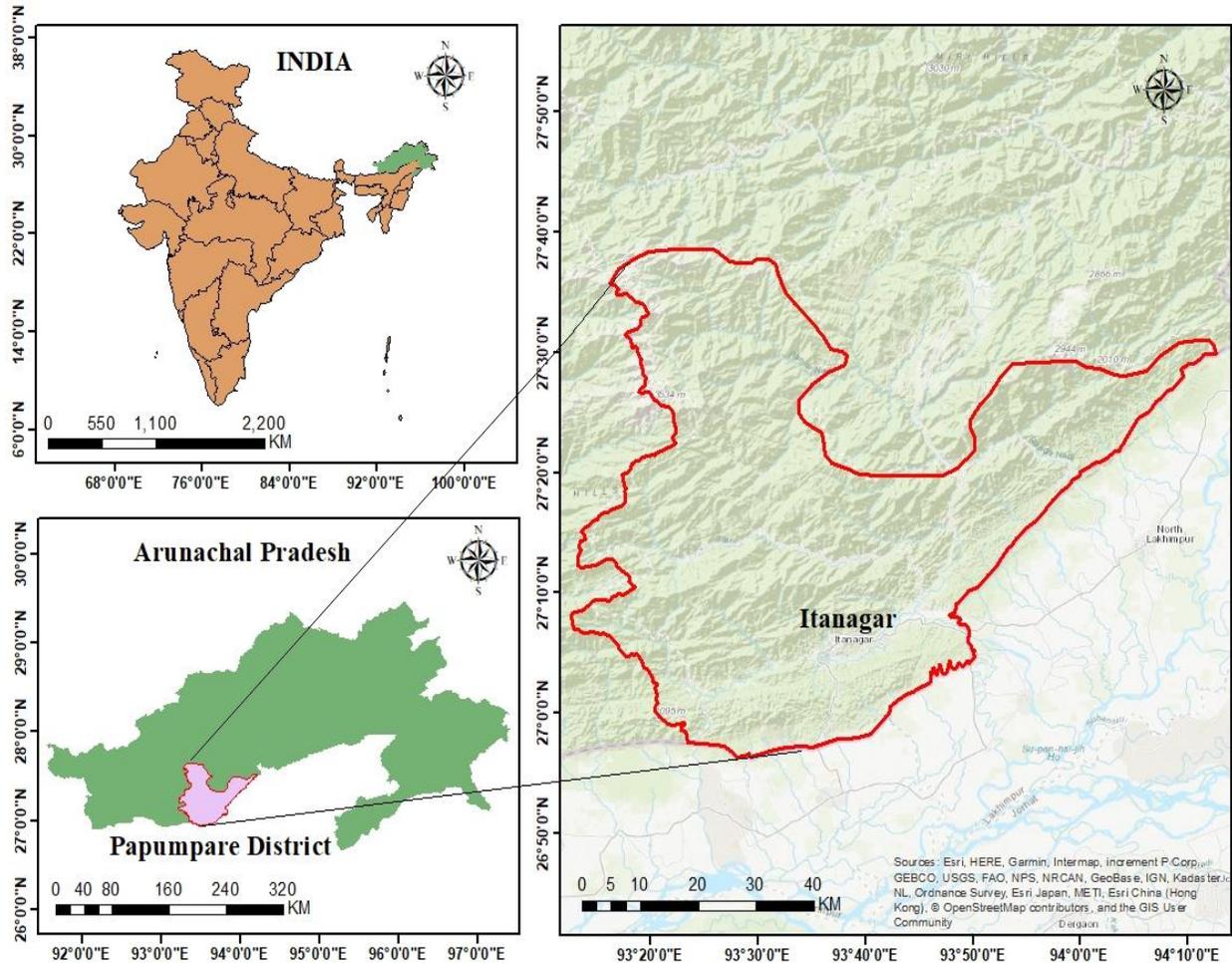
144 conditions along with frequency and intensity of potential earthquakes. This region consist of
145 highly variable topography characterized by hilly terrain and river valleys, leads significant effects
146 on seismic attenuation, which could exacerbate the impact of seismic events on structures and
147 infrastructure. By employing PSHA framework, which integrates the inherent uncertainties
148 associated with earthquake occurrence, ground motion prediction, and site effects, this study
149 enables a more robust estimate of seismic hazard, providing decision-makers with critical
150 information to develop mitigation strategies. For a rapidly developing city like Itanagar, where
151 infrastructure projects are underway, the results of a PSHA are crucial for designing buildings and
152 public facilities that can withstand seismic forces, thus minimizing casualties and economic losses
153 in the event of a significant earthquake. The PSHA model for Itanagar will involve the evaluation
154 of different seismic source zones, ground motion attenuation relationships, and local soil
155 characteristics. This approach provides seismic hazard curves, maps of PHA, and spectral
156 acceleration values for different return periods that will help planners and engineers to design and
157 construct the seismically resilient structures. The present study focuses on the assessment of
158 seismic hazard for Itanagar region in the state of Arunachal Pradesh (India) in view of planning a
159 proposed smart city. Moreover, PHA and spectral acceleration (S_a) are used to formulate the
160 outcomes of the seismic hazard analysis at various return periods. PHA provides insight to ground
161 deformation, strain development, horizontal forces, and shear stresses, which are essential for
162 earthquake-resistant design considerations; whereas, S_a represents the maximum acceleration
163 experienced by structures represented by a single-degree-of-freedom system under damped
164 vibrations [39]. The prepared hazard maps of Itanagar region corresponding to 475 years (10%
165 probability of exceedance in 50 years) and 2475 years (2% probability of exceedance in 50 years)
166 at different time periods, such as 0.1 s, 0.3 s, 0.5 s, 1.0 s, 2.0 s, and 3.0 s, will help engineers and
167 architects to plan and design of earthquake resistant structures on the highly undulated hilly terrain.

168

169 **2. Study area and tectonic features**

170 Itanagar is located in Arunachal Pradesh, which lies on the northeastern region of India, at
171 27°05'54"N and 93°37'19"E, as shown in Fig. 1. The entire state of Arunachal Pradesh covers an
172 area of 83743 km², and is located in the Siwalik range of Himalaya, encompassing a range of
173 elevations from 102 m to 588 m above mean sea level (MSL) [40]. It is a physiographic section of
174 the great Himalayas, wherein Lohit, Dibang, Siang, Kameng, and Subansiri rivers are the most

175 influencing ones among many other rivers and their tributaries. In the Himalayan Fold Thrust
 176 (HFT) belt, Itanagar lies within the active seismic domain near the plate boundary and falls under
 177 Zone-V, the highest seismic vulnerability category as per IS 1893-Part-1 [32].



178
 179 Fig. 1 Location of Itanagar in Arunachal Pradesh, India
 180

181 The entire Itanagar consist of mountain slope faces, ridges and crest lines, open slopes, and mid-
 182 slope ridges with spatial extents for buildings, roads, drainage, and sewage networks. The region
 183 also consists of active seismotectonic domain surrounding the Itanagar urban agglomeration zone
 184 [41-43]. This zone is a result of collision between Indian with Eurasian plate [44,45]. Main
 185 boundary thrusts (MBT) and main central thrusts (MCT) are most likely to cause seismic events,
 186 wherein MCT reflects ductile shear zones [46]. A transverse tectonic regime throughout NE
 187 Himalayan belt is also observed with a focus depth ranging from 0 to 70 km [46]. It is reported that
 188 there were two earthquakes of magnitude 7.1 and 7.8 in 1941 and 1947 in this region. The 2011

189 Sikkim earthquake of M_w 6.8 caused severe devastation in the Sikkim region of Northeast India.
190 The Shillong plateau, with focal depth ranging from 0-60 km, is bordered by the Brahmaputra river
191 fault in the north and the Dauki fault in the south. The west side consists of Dhubri fault that is
192 oriented north-south, while the east side has the Disang thrust. Moreover, the Sylhet fault is
193 responsible for 1918 Srimangal earthquake ($M_w = 7.6$). Shillong plateau is much more seismically
194 active than Naga thrust zone. It is also called as Assam gap or aseismic corridor by the researchers
195 [47,48]. This gap runs parallel to the Dauki fault in the south and extends parallelly to the Naga
196 Thrust in the east. Due to the collision between Indian and Eurasian plates in Mishmi Thrust zone,
197 seismicity of the entire region becomes significantly higher than the Eastern Himalaya [49]. As a
198 result of the high stress concentration, this zone is classified as a special zone with block tectonics
199 [50]. To quantify the impact of these tectonic processes and enhance our understanding of seismic
200 patterns, the next step involves the systematic development and refinement of an earthquake
201 catalogue for the Itanagar region.

202

203 **3. Development of earthquake catalogue, homogenization and declustering**

204 To estimate the hazards associated with earthquakes, a comprehensive earthquake catalogue is
205 necessary. The data statistics of seismic events are essential to assess the seismic hazard of a region.
206 The earthquake catalogue used in this study covers the earthquake period from 1900 to 2024,
207 collected from United States Geological Survey (USGS), National Centre of Seismology (NCS)
208 under the Government of India and International Seismological Centre (ISC). An earthquake
209 catalogue is developed using seismic events recorded within a 500 km radius around Itanagar and
210 includes approximately 2710 earthquake events of magnitudes exceeding 4.0 i.e., ranging from a
211 minimum moment magnitude of 4.4 to a maximum of 8.0. This radius is selected based on
212 established practices in seismic hazard studies for regions with complex tectonic environments
213 and high seismic activity. Previous seismic hazard studies in Northeast India have effectively used
214 similar radii [35, 36, 51] to adequately capture significant seismic sources influencing regional
215 seismic hazards.

216

217 A complete and consistent earthquake catalogue is crucial for understanding seismic activity in a
218 region. The regional earthquake catalogues are often heterogeneous due to variations in magnitude
219 scales. Therefore, to ensure uniformity in the completeness analysis, this study converted different

220 magnitude scales into a moment magnitude using the correlations proposed by Scordilis [52], as
 221 described in Eqns. 1-3. This standardization helps address the saturation issues associated with
 222 various magnitude scales. The catalogue includes essential information such as event coordinates,
 223 date (month and year), magnitude, and hypocentral depth.

$$224 \quad M_w = 0.85 M_b + 1.03, \text{ for } 3.5 \leq M_b \leq 6.2 \quad (1)$$

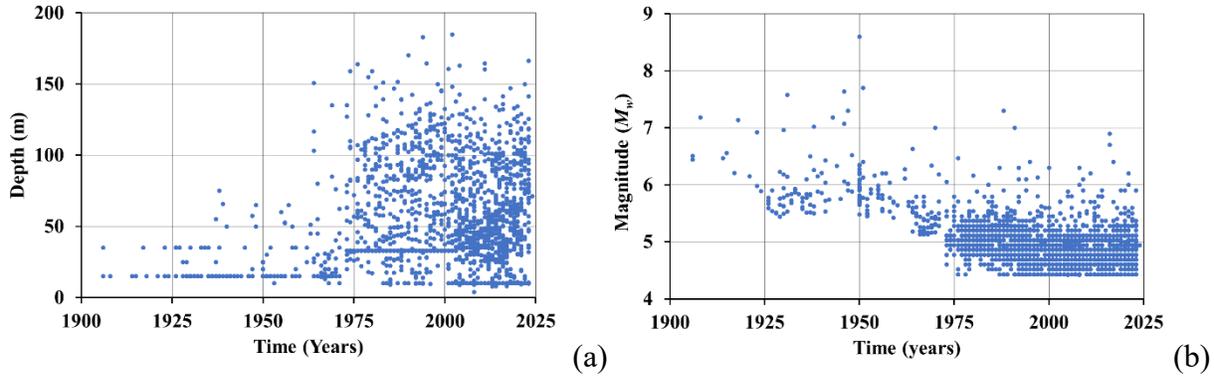
$$225 \quad M_w = 0.67 M_s + 2.07, \text{ for } 3.0 \leq M_s \leq 6.1 \quad (2)$$

$$226 \quad M_w = 0.99 M_s + 0.08, \text{ for } 6.2 \leq M_s \leq 8.2 \quad (3)$$

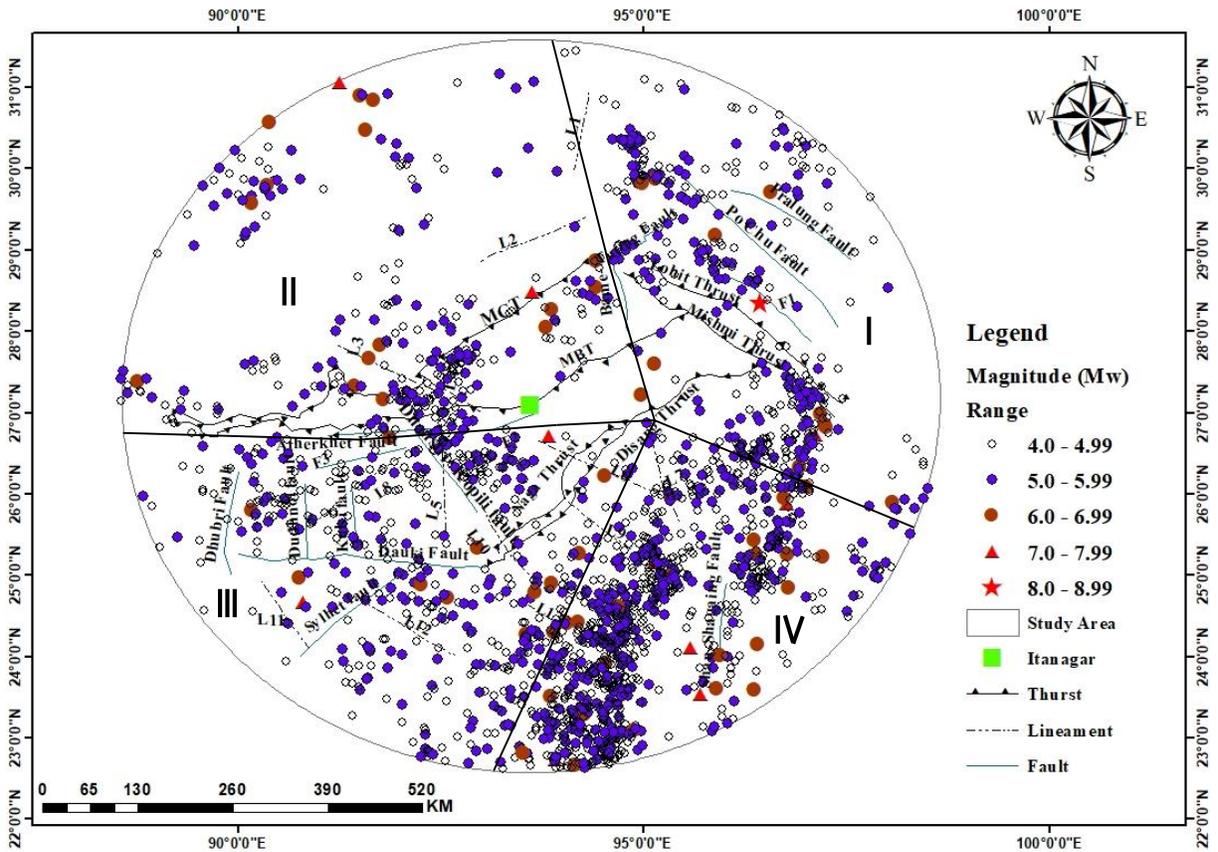
227
 228 To convert local magnitude (M_L) to moment magnitude (M_w), regional correlations are typically
 229 preferred for better accuracy. For India and its surrounding regions, the correlation derived by
 230 Kolathayar *et al.* [53] is utilized, as expressed in Eqn 4.

$$231 \quad M_w = 0.815 M_L + 0.767, \text{ for } 3.3 \leq M_L \leq 7.0 \quad (4)$$

232
 233 The expressions for magnitude conversion provided by Scordilis [52] and Kolathayar *et al.* [53]
 234 are chosen based on their extensive validation for tectonically active regions and their relevance
 235 to the Himalayan and Indo-Burmese seismic zones, where Itanagar, Arunachal Pradesh, is located.
 236 Further, declustering of the homogenized catalogue is conducted to better estimate the earthquake
 237 return periods by removing foreshocks and aftershocks. Over the years, various methods have been
 238 developed globally for declustering by researchers [54–57]. In this study, declustering is conducted
 239 using a homogenized earthquake catalogue in ZMAP v2007 [58], following the algorithm
 240 proposed by Gardner and Knopoff [54]. Figure 2a illustrates the variation in seismic event depths
 241 over time, showing focal depth ranging between 10 and 120 km for most significant earthquakes
 242 in this region. Figure 2b depicts the magnitudes of seismic events over time, revealing that
 243 earthquakes with magnitudes of $M_w \leq 5.0$ were only recorded after 1964, thereby hinting the
 244 absence of monitoring stations in the region prior to that period. From the initial dataset of 2710
 245 events, 24% were identified as dependent and removed during declustering; thus, a total of 2054
 246 mainshock event dataset is considered for further analysis. The seismic activity in the study region
 247 includes 1184 mainshocks with magnitudes between 4.0–4.9 M_w , 793 between 5.0–5.9 M_w , 64
 248 between 6.0–6.9 M_w , 12 between 7.0–7.9 M_w , and one event is 8.0 M_w . Figure 3 presents the
 249 epicenter map of the study area, displaying all 2054 earthquake events that is included in the final
 250 catalogue.



251
 252 Fig. 2 Temporal distribution of declustered seismic events: (a) focal depths and (b) magnitudes,
 253 recorded between 1906 and 2024 within the influence zone of study area
 254



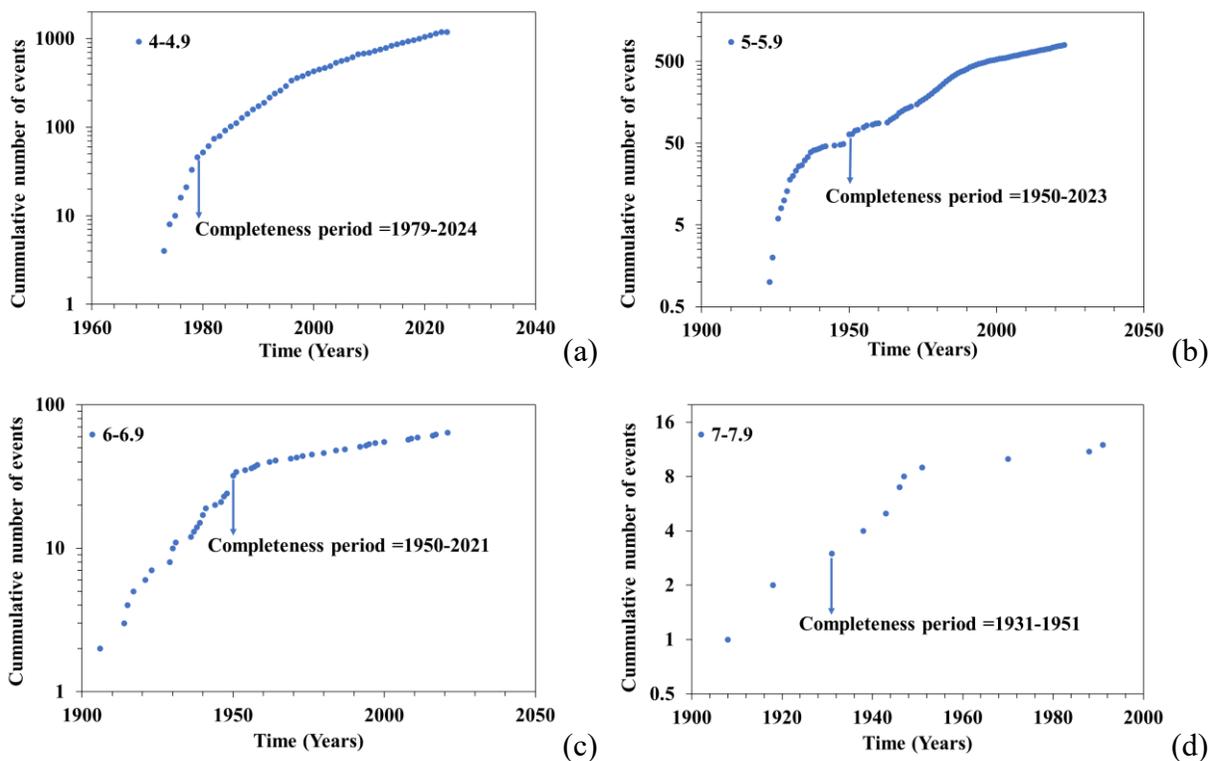
255
 256 Fig. 3 Seismicity map of study area
 257

258 **4. Completeness of data with respect to magnitude and time**

259 The number of earthquakes per decade is categorized into five magnitude ranges: $4 \leq M_w \leq 4.9$, 5
 260 $\leq M_w \leq 5.9$, $6 \leq M_w \leq 6.9$, $7 \leq M_w \leq 7.9$, and $8 \leq M_w \leq 8.9$. Table 1 provides a detailed count of

261 earthquakes recorded in each decade, starting from the earliest available earthquake data. Figure 4
 262 presents a histogram illustrating the data in Table 1 for the entire catalogue, spanning from 1906
 263 to 2024. Developing an earthquake catalogue that encompasses a significant time period is crucial,
 264 as incomplete catalogues can lead to inaccurate estimation of seismicity parameters. To ensure
 265 completeness, the catalogue was analyzed using the Cumulative Visual Inspection (CUVI) method
 266 [59]. This method assesses completeness by plotting the cumulative number of events per year
 267 against the time of occurrence for each magnitude range. The completeness period is determined
 268 as the year from which there is a noticeable steep rise in the graph. Based on the CUVI method,
 269 the catalogue is deemed complete for large magnitude earthquakes throughout the entire time span.
 270 Table 2 demonstrates that the catalogue achieves completeness over a sufficient duration, making
 271 it reliable for seismic analysis.

272



274

275

276 Fig. 4 Catalogue completeness by the CUVI method for the study area for various classes of
 277 earthquake magnitude

278

279

280

Table 1 Number of earthquakes reported in each decade for study region

Years Interval	Number of Earthquake					Total
	4.0-4.9	5.0-5.9	6.0-6.9	7.0-7.9	8.0-8.9	
1906-1915			4	1		5
1916-1925		2	3	1		6
1926-1935		29	4	1		34
1936-1945		16	9	2		27
1946-1955		31	15	4	1	51
1956-1965		24	6			30
1966-1975	10	67	3	1		81
1976-1985	92	154	4			250
1986-1995	189	155	5	2		351
1996-2005	267	98	2			367
2006-2015	307	110	4			421
2016-2024	319	107	5			431
Total	1184	793	64	12	1	2054

281

282

283

Table 2 Completeness analysis of earthquake catalogue

Magnitude class (M_w)	Completeness Analysis (CUVI method)	
	Period	Interval (Years)
4.0-4.9	1979-2024	45
5.0-5.9	1950-2023	73
6.0-6.9	1950-2021	71
7.0-7.9	1931-1951	20

284

285 5. Exploration and characterization of potential seismogenic sources

286 The process of characterizing earthquake sources includes creating tectonic maps, identifying all
 287 potential sources of damaging earthquakes, assessing the largest recorded earthquake magnitudes,
 288 measuring the lengths of faults, lineaments, and thrusts, and estimating the maximum magnitude
 289 that these seismogenic sources can generate. Measuring fault lengths is significant because it helps
 290 in understanding the size and location of potential earthquakes. By determining the dimensions of
 291 faults, researcher can assess seismic activity potential and better prepare for future earthquakes.
 292 Additionally, this information aids in assessing risk to populated areas and implementing safety
 293 measures. Estimating the magnitude potential (M_{max}) of seismogenic sources is crucial for
 294 assessing the maximum magnitude of earthquakes that can occur in a given region.

295 **5.1 Identification of seismogenic sources and development of tectonic setup**

296 The Geological Survey of India (GSI) developed a Seismo-tectonic Atlas of India and its Environs
297 (SEISAT) as a detailed resource for analyzing seismic activity in the area [60]. This atlas presents
298 tectonic features and earthquake epicenters on a 1:1000000 scale across 43 sheets, each
299 encompassing 3° longitudes and 4° latitudes. For this research, a 500 km radius surrounding
300 Itanagar is examined using high-resolution scans of sheets 13-17. A tectonic map of the seismic
301 study area was created by digitizing and combining these sheets, as illustrated in Fig. 4. SEISAT
302 categorizes all linear tectonic features as either active or inactive. The distribution of seismic
303 events across various tectonic features is revealed by overlaying recorded event epicenters on
304 tectonic maps. This analysis showed that the seismic events occurred most frequently in zones
305 with active tectonic features, providing valuable insights for understanding of seismic hazards in
306 the region. In seismically active areas, there is no standardized method for identifying potential
307 seismogenic sources [61]. Consequently, tectonic maps and historical earthquake epicenter
308 locations are typically used as primary references in these assessments. Potential seismogenic
309 sources are identified based on maximum observed magnitude (M_{obs}) values and their proximity
310 to the site of interest for seismic hazard assessment. The seismic study region contains 33 active
311 tectonic features, as shown in Fig. 4. This study considers 18 major active tectonic features capable
312 of producing significant ground motion at the selected site for seismic hazard analysis, as listed in
313 Table 3.

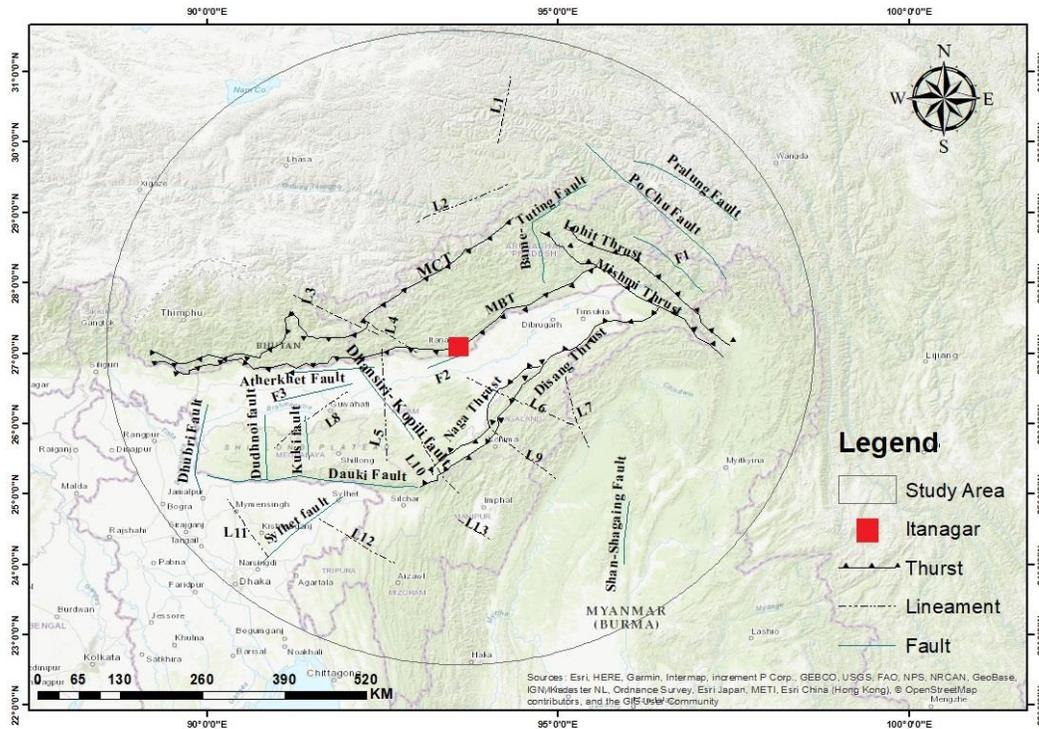


Fig. 5 Seismotectonic map showing tectonic features of the study region

5.2 Estimation of maximum magnitude potential

The maximum potential earthquake magnitude (M_{max}) plays a crucial role in seismic hazard evaluation and the earthquake-resistant design of structures. It represents the largest seismic event in an area that can be produced at a particular source. Accurate estimation of M_{max} is essential for developing effective earthquake risk reduction plans, as it directly affects the design criteria for buildings, bridges, and other vital infrastructure to endure possible seismic occurrences. Several researchers have proposed various techniques to estimate M_{max} over the time, acknowledging its significance in minimizing earthquake-related risks [62-66]. These approaches differ in their methodologies, reflecting the intricacy and variability of seismic sources across different regions. This research concentrates on the methods suggested by Gupta [63] and National Disaster Management Authority (NDMA) [66] for estimating M_{max} . Gupta [63] presented a general approach that involves adding 0.5 units to the maximum observed magnitude (M_{obs}) to determine M_{max} . This method is founded on the principle of historically maximum observed seismic magnitude that provides a baseline, which necessitates a conservative adjustment to the account of uncertainties and variations in seismic activity. In contrast, NDMA [66] approach offers a slightly

332 different method, adjusting M_{max} on the observed magnitude. Specifically, if M_{obs} is below 5.0, a
333 smaller increment of 0.3 is added, indicating a lower likelihood of significant magnitude escalation
334 for smaller seismic events. Conversely, if M_{obs} is 5.0 or higher, a larger increment of 0.5 is applied,
335 recognizing the increased probability of more substantial seismic activity in areas that have already
336 experienced moderate to strong earthquakes. These methodologies are widely accepted and
337 utilized because they provide a standardized, yet adaptable, approach to estimating M_{max} , ensuring
338 that seismic hazard assessments are both realistic and conservative. These conventional and
339 broadly adopted methods, which involve adding constant incremental values to M_{obs} , are
340 particularly valuable for seismic sources with limited historical data, as they offer a systematic
341 way to estimate potential maximum magnitudes in the absence of comprehensive long-term
342 records. The application of these methods is supported by numerous studies in the field [11, 37,
343 67], which have demonstrated the effectiveness of these approaches in various seismically active
344 regions.

345

346 Table 3 Estimation of M_{max} for various existing faults in the influence zone of study area

Fault Name	Total Length	M_w (observed)	Estimated M_{max}	
			Gupta [63]	NDMA [66]
MCT	700	7.3	7.8	7.8
Lohit Thurst	326	8.0	8.5	8.5
Mishmi Thurst	354	6.3	6.8	6.8
MBT	737	7.3	7.8	7.8
Naga Thurst	267	5.5	6.1	6.1
Shan-Shagaing Fault	106	7.6	8.1	8.1
Bame-Tuting Fault	209	6.4	6.9	6.9
Dhubri Fault	148	7.1	7.6	7.6
Dhansiri Kopili Fault	141	6.0	6.5	6.5
Atherkhet Fault	133	6.0	6.5	6.5
Dudhnoi Fault	106	5.5	6.1	6.1
Dauki Fault	319	7.1	7.6	7.6
Sylhet Fault	166	6.7	7.2	7.2
Pralung Fault	186	6.1	6.6	6.6
Po-Chu Fault	308	6.3	6.8	6.8
F2	60	4.9	5.4	5.2
L5	185	4.8	5.3	5.1
L6	178	6.5	7.0	7.0

347 **6. Gutenberg- Richter seismicity parameters**

348 The seismic study region is segmented into four distinct regions, encompassing 18 tectonic
349 features. To estimate seismicity parameters for the region, the average observed focal depths within
350 each seismogenic area are utilized [67]. Based on the developed earthquake catalog, the study area
351 is divided into four regions: Region-I, Region-II, Region-III, and Region-IV, as shown in Fig. 3.
352 According to Anbazhagan *et al.* [68], several recurrence laws describe the variability in earthquake
353 magnitudes generated by different seismic sources, including the Gutenberg-Richter (G-R) relation
354 [69] and the Mertz-Cornell model [70]. Among these, the G-R relation is simple and widely
355 employed for evaluating the seismic hazard parameter, ‘*b*’. For conducting a Probabilistic Seismic
356 Hazard Analysis (PSHA), the recurrence parameters ‘*a*’ and ‘*b*’ are critical. These parameters can
357 be determined using the G-R recurrence law, which assumes that earthquakes occur in any given
358 region following a Poisson distribution, implying independence in their timing and location. The
359 Gutenberg-Richter law is expressed mathematically as follows:

$$360 \log \lambda_m = a - bM_w \quad (5)$$

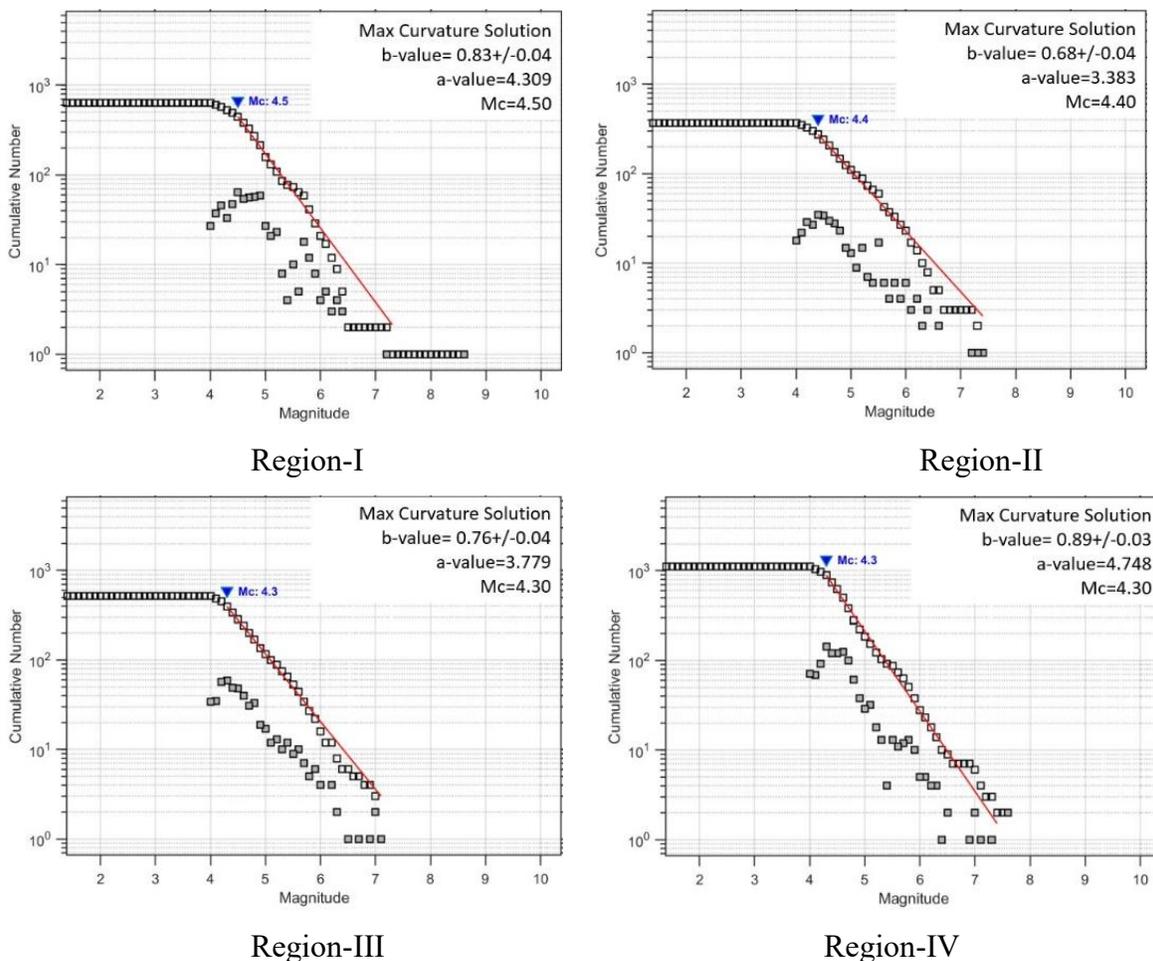
361
362 In the Gutenberg-Richter relationship, the parameters ‘*a*’ and ‘*b*’ describe the seismicity of a
363 region, while λ_m represents the average rate of exceedance for a moment magnitude M_w . The
364 parameter ‘*a*’ (intercept) is the logarithm of the total number of earthquakes with magnitudes equal
365 to or exceeding the threshold magnitude. The parameter ‘*b*’ (slope) reflects the average distribution
366 of earthquake magnitudes in a specific area. A lower *b*-value indicates a predominance of larger
367 magnitude earthquakes, while a higher *b*-value suggests that smaller magnitude earthquakes are
368 more common. According to NDMA [66], the value of ‘*b*’ varies from region to region, it lies
369 typically in the range $0.6 < b < 1.5$. The *b*-value of the present study ranges between 0.68 ± 0.04 to
370 0.89 ± 0.03 . The Gutenberg-Richter relationship for the present study area is shown in Fig. 6.
371 Regions with lower *b*-values (e.g., Region II) are typically more prone to significant seismic
372 hazards due to the higher frequency of large earthquakes. Regions with higher *b*-values (e.g.,
373 Region IV) are more stable with less risk of catastrophic seismic activity but experience more
374 frequent smaller earthquakes.

375
376 The seismicity parameters (*a* and *b* values) estimated for the Itanagar region are summarized in
377 Table 5. The *b*-value, representing the relative frequency distribution of earthquake magnitudes, is

378 found to range from 0.64 to 0.92, aligning closely with previously reported values for seismically
 379 active regions in Northeast India, such as 0.43–1.07 by Das *et al.* [35] and 0.54–0.86 by Sil *et al.*
 380 [36]. This similarity suggests that moderate-magnitude earthquakes dominate the seismic activity
 381 in Itanagar, consistent with the regional tectonic framework. The *a*-value, which characterizes the
 382 overall seismic activity and event rate, ranges from 3.77 to 4.75 in the present study. These values
 383 correspond well to those reported from similar tectonic and seismic settings, including 4.21 for
 384 Peninsular India (Jaiswal and Sinha [71]), 3.52 for Bangalore (Anbazhagan *et al.* [27]), and a range
 385 of 2.54–4.94 for Tripura and Mizoram (Sil *et al.* [36]). Generally, higher *a*-values indicate regions
 386 with greater seismic event frequencies, thus underscoring higher overall seismic activity. The
 387 obtained *a*-values for Itanagar clearly reflect a significant level of seismic activity, confirming the
 388 need for detailed seismic hazard assessments and earthquake-resistant infrastructure design in this
 389 area.

390

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394

Fig. 6 Frequency-magnitude relationships (Gutenberg–Richter law) for four distinct regions.

395

Table 4. Seismicity parameters for different area sources

Region	b	a	Number of events	Range of Magnitude	Maximum M_{obs}	Return period for M_{obs} (in Years)
I	0.83±0.04	4.309	427	4.0-8.0	8.5	674
II	0.68±0.04	3.383	271	4.0-7.4	7.4	45
III	0.76±0.04	3.779	412	4.0-7.1	7.1	42
IV	0.89±0.03	4.748	923	4.0-7.6	7.6	104

396

397

Table 5. Comparison of ‘ b ’ values obtained from present study with the previous literature.

Authors	Study area	b -value	a -value	Data period (years)
Ram and Rathor [70]	South India	0.81	-	70
Kaila <i>et al.</i> [73]	South India	0.7	-	14
Rao and Rao [72]	Peninsular India	0.85	4.4	170
Raghukanth and Iyenger [75]	Mumbai	0.86	0.77	-
Jaiswal and Sinha [71]	Peninsular India	0.92	4.21	160
Anbazhagan <i>et al.</i> [27]	Bangalore	0.86	3.52	200
Vipin <i>et al.</i> [76]	South India	0.891	4.58	400
Menon <i>et al.</i> [77]	Tamil Nadu	1.13	5.05	501
Sitharam <i>et al.</i> [78]	Karnataka	0.923	4.75	400
Shukla and Choudhury [79]	Gujarat	0.51	-	188
Kolathayar <i>et al.</i> [23]	India	0.5-1.5	3-10	1760
Kumar <i>et al.</i> [80]	Lucknow	0.80-0.86	3.2-4.07	170
Sil <i>et al.</i> [36]	Tripura & Mizoram	0.54-0.86	2.54-4.94	279
Naik <i>et al.</i> [81]	Goa	0.91	6.41	246
Shiuly <i>et al.</i> [82]	Kolkata	0.738	2.73	120
Das <i>et al.</i> [35]	Northeast region	0.43-1.07	1.68-5.76	113
Present study	Itanagar	0.64-0.92	3.77-4.7	124

398

399 7. Probabilistic seismic hazard analysis

400 The hazard analysis in this study was conducted using R-Crisis v18.2 that is a windows-based
 401 software developed by Ordaz and Salgado-Gálvez [83]. R-Crisis is specifically designed for PSHA
 402 and calculates seismic hazard by incorporating earthquake occurrence probabilities, attenuation
 403 patterns, and seismicity trends. For PSHA, the software supports three types of probable source
 404 geometries: areas, lines, and points. The analysis utilizes the earthquake catalogue for the study
 405 region in a homogenized form of the moment magnitude scale. Key seismic parameters such as
 406 maximum magnitude (M_{max}), earthquake activity rate (k), and the b -value are calculated using the
 407 frequency–magnitude distribution proposed by Gutenberg and Richter [69]. One of the critical
 408 outcomes of the PSHA is the estimation of PHA, which quantifies the intensity of ground shaking

409 during an earthquake at a specific location. This study presents a summary of PHA values for the
410 Itanagar region for various structural periods and return periods, providing valuable insights into
411 the seismic risk for the area.

412
413 Consequently, the development of a new Ground Motion Prediction Equation (GMPE) is not
414 undertaken. Globally available GMPEs are reviewed to identify a suitable model for the Itanagar
415 area, which is located in a highly seismically active region and is particularly vulnerable to shallow
416 crustal earthquakes, similar to those in the Northeast Himalayan region [43, 84]. To select an
417 appropriate GMPE, several models for shallow crustal earthquakes are evaluated by comparing
418 the observed and calculated PGA values for past seismic events, including the 6.8 M_w Sikkim
419 earthquake of 2011 and other regional events [85]. Following this analysis, the attenuation model
420 by Boore *et al.* [86] is identified as the best-fit model for the study area, based on its superior
421 agreement with observed seismic data. Due to limited and publicly unavailable earthquake data, it
422 becomes even harder to deal with different types of uncertainties [87]. As a result, global models
423 like NGA-West2 [84] are often used for seismic hazard studies in India and the Himalayan region.
424 Previous studies, such as Sharma *et al.* [88] and Ornthammarath *et al.* [89], have demonstrated the
425 reliability of the Boore *et al.* [84] GMPE in seismic hazard analyses for Himalayan and Indo-
426 Burmese regions, where established local GMPEs are currently unavailable. Additionally, Ghione
427 *et al.* [90] and Lallawmawma *et al.* [91] support the application of global GMPEs in seismic hazard
428 assessments for Northeast India, emphasizing their suitability in estimating strong ground motions.
429 The selection of the Boore *et al.* [86] model in this study is further justified by the similarities in
430 tectonic stress regimes, seismogenic depths, and geological conditions between the study area and
431 other seismically active areas where this GMPE has been validated. While the importance of
432 developing region-specific GMPEs is acknowledged, the current absence of adequate strong-
433 motion data from study region necessitates the adoption of a widely tested, globally validated
434 GMPE, ensuring robust and reliable ground motion estimates. The seismic hazard is assessed for
435 return periods of 475 years (10% probability of exceedance in 50 years) and 2475 years (2%
436 probability of exceedance in 50 years). Using R-Crisis, which supports various verified GMPEs,
437 the analysis generated exceedance probability plots and stochastic event simulations. The input
438 parameters for the computational framework included seismicity constants (a , b), minimum and
439 maximum magnitudes (M_{\min} , M_{\max}), focal depths, and the selected attenuation model. In this study,

440 these parameters include the precise hypocentral depths, which are determined based on historical
441 earthquake records, with average depths calculated as 35 km, 28 km, 40 km and 70 km for Regions
442 I, II, III and IV, respectively. The software also requires the information about fault geometry and
443 characteristics of seismic sources, for which 18 major tectonic faults (identified through the
444 Seismotectonic Atlas SEISAT, GSI, and digitized in ArcGIS software) are modeled as line
445 sources. The PHA values for the Itanagar region are calculated across a grid of points, and the
446 results are represented as contour maps. For a return period of 475 years, the estimated PHA is
447 0.22g, corresponding to a 10% probability of exceedance in a 50year period (Fig. 7a). For a
448 2475year return period, the PHA increases to 0.36g, representing a 2% probability of exceedance
449 in the same timeframe (Fig. 7b). These contour maps provide a visual representation of the seismic
450 hazard distribution in the region.

451

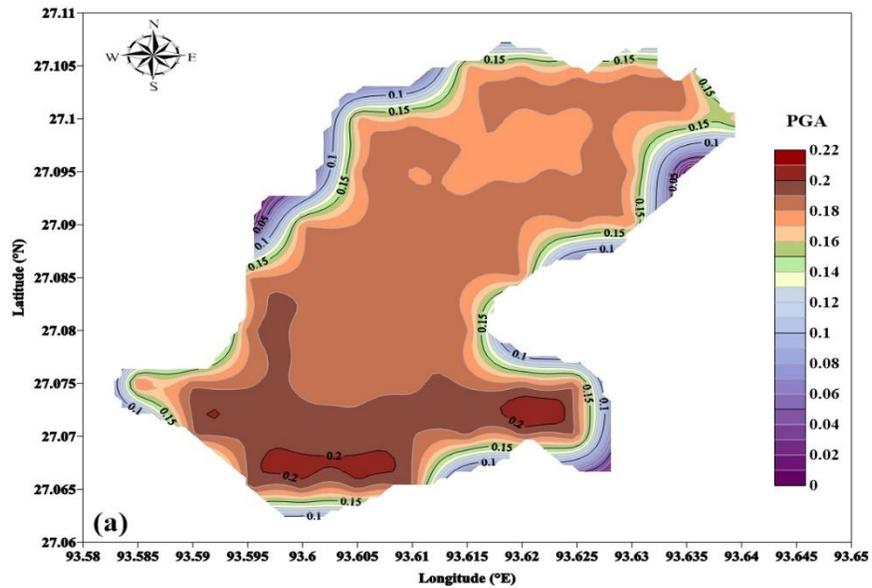
452 It is noteworthy that the Indian code of practice, IS 1893-Part-1 (2016) [32], classifies the entire
453 northeastern region of India, including the study area, as seismic zone V, assigning a PHA value
454 of 0.36g without referencing a specific return period. In this study, the computed PHA values for
455 return periods of 475 years and 2475 years are found to be consistent with the value recommended
456 by IS: 1893. Specifically, the PHA values of 0.22g for a 475 year return period and 0.36g for a
457 2475year return period align well within the prescribed limit of 0.36g, supporting the reliability of
458 the hazard estimates. The seismic hazard assessment for the Itanagar region, based on spectral
459 acceleration (S_a) contours at a 475year return period across periods of 0.1 s, 0.3 s, 0.5 s, 1.0 s, 2.0
460 s and 3.0 s, reveals significant spatial variability in seismic intensity, have been plotted as contour
461 map in Fig. 8(a-f).

462

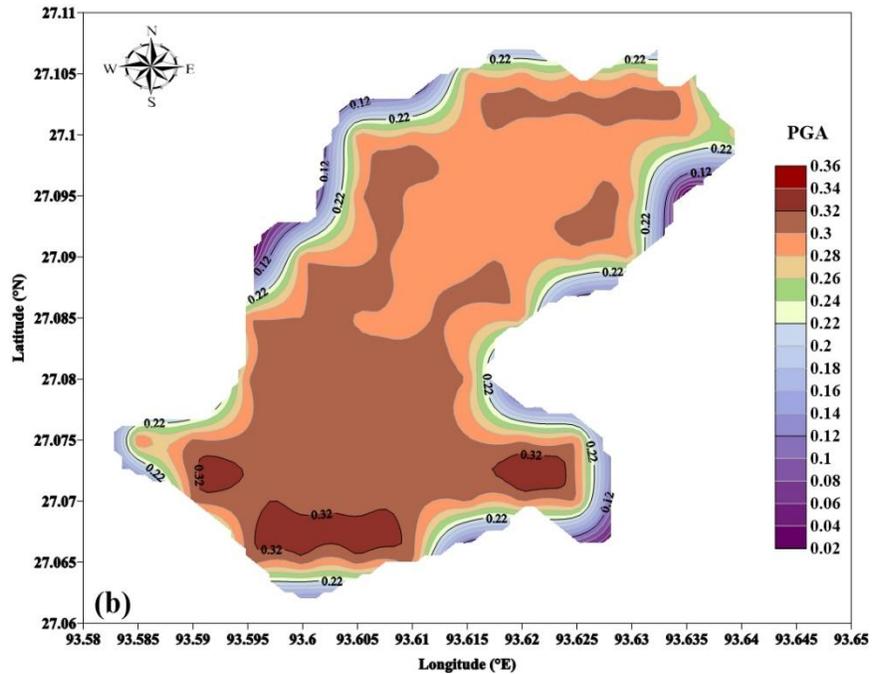
463 The southern and central areas consistently exhibit higher S_a values, peaking at 0.54g for the 0.1 s
464 period, indicating elevated seismic hazard. In contrast, the northern and western regions display
465 lower S_a values, with a minimum of 0.024g at the 3.0 s period, suggesting reduced seismic risk.
466 These spectral acceleration values indicate the potential severity of ground shaking during an
467 earthquake, with higher values suggesting more intense shaking. At 0.54g for 0.1 s, the risk of
468 structural damage is significant, especially for buildings not designed to withstand such forces. As
469 the spectral acceleration decreases with longer periods, the risk for taller structures or those with

470 longer natural periods may still be considerable if they are not adequately engineered for these
471 conditions.

472



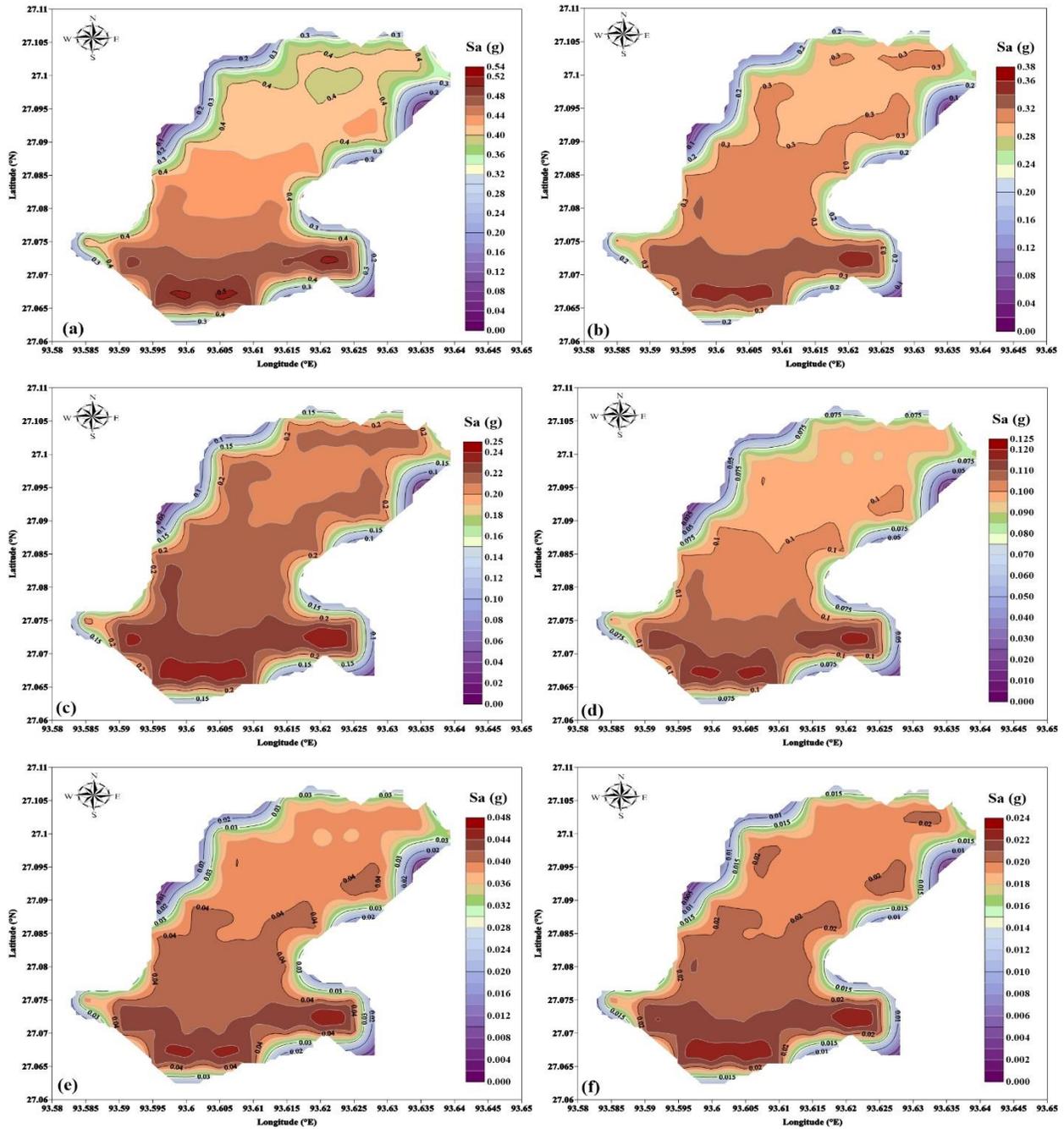
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475 Fig. 7 PHA map of Itanagar for (a) 10% probability of exceedance in 50 years for return period
476 of 475 years (b) 2% probability of exceedance in 50 years for return period of 2475 years

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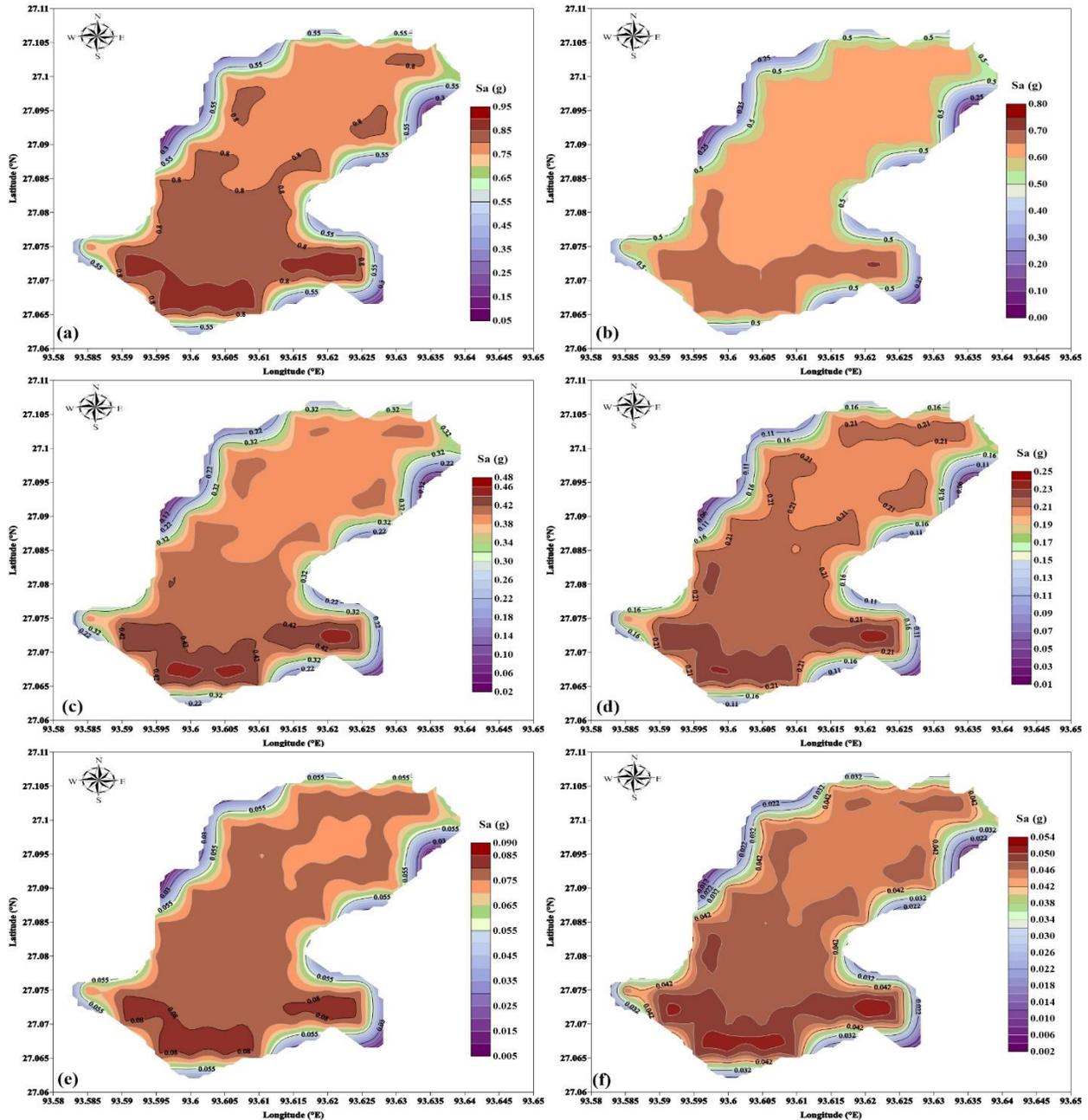
481 Fig. 8 Spectral acceleration in g at periods of (a) 0.1 s (b) 0.3 s (c) 0.5 s (d) 1.0 s (e) 2.0 s
 482 (f) 3.0 s, for a 475 year return period

483

484 The spectral acceleration contours for the study region, at return periods of 2475 year, at periods
 485 of 0.1 s, 0.3 s, 0.5 s, 1.0 s, 2.0 s and 3.0 s, is presented as contour plots in Fig. 9(a-f). The results
 486 indicate that the southern and central regions exhibit consistently higher S_a values across all
 487 periods, suggesting these areas are more susceptible to severe ground motion during earthquakes.

488 At the shortest spectral period (0.1 s), S_a values reach a maximum of 0.95g in the southern region,
 489 reflecting high seismic intensity likely due to local soil amplification and tectonic activity.
 490 Conversely, the northern and western regions demonstrate relatively lower S_a values, with ranges
 491 of 0.20g to 0.35g at 0.1 s, gradually decreasing to 0.02g to 0.04g at 3.0 s, indicating lower seismic
 492 hazard in these areas.

493



494

495

496

497 Fig. 9 Spectral acceleration in g at periods of (a) 0.1 s (b) 0.3 s (c) 0.5 s (d) 1.0 s (e) 2.0 s

498

(f) 3.0 s, for a 2475 year return period

499 Table 6. Comparison of PHA values from the present study with previous studies for Northeast
 500 India

References	PHA (g) for 10% in 50 years exceedance
Bhatia <i>et al.</i> [93]	0.25-0.45
Das <i>et al.</i> [92]	0.18–0.22
IS-1893 [38]	0.18
Sharma and Malik [94]	0.3-0.48
Desai and Choudhury [31]	0.095-0.2
NDMA [66]	0.15-0.35
Nath and Thingbaijam [95]	0.5-1.12
Pallav <i>et al.</i> [51]	0.13-0.19
Borgohain <i>et al.</i> [38]	0.59
Bahuguna and Sil [37]	0.2-0.48
Kumar <i>et al.</i> [96]	0.25
Shukla and Choudhury [28]	0.23-0.34
Present Study	0.22

501
 502 As the spectral period increases, S_a values decrease across all regions, with the highest values at
 503 3.0 s remaining in the central and southern areas, peaking at 0.08g-0.09g. The central region
 504 consistently exhibits moderate to high S_a values across all periods, highlighting it as a critical area
 505 for seismic design considerations. The seismic hazard map for both the return period shows
 506 uniform hazard inside the zones and sharp changes in hazard values at their edges. The observed
 507 decrease in S_a values with increasing spectral periods reflects the attenuation of seismic energy
 508 over time, a trend consistent with established seismic hazard principles. This analysis underscores
 509 the importance of incorporating spatially resolved spectral acceleration data into the seismic design
 510 of infrastructure and urban planning to mitigate earthquake risks effectively. The findings of the
 511 current study are compared with the results of other researchers, focusing on the return period of
 512 475 years. A detailed comparison is presented in Table 6, highlighting the similarities and
 513 differences in PHA values. The present study reports a PHA value of 0.22g for 10% probability of
 514 exceedance in 50 years for return period of 475 years, which aligns closely with findings from
 515 similar studies. Das *et al.* [92] provides a range of 0.18–0.22g, overlapping entirely with the present
 516 study, indicating consistent seismic hazard levels. IS-1893 specifies a value of 0.18g, which is
 517 slightly lower than the values obtained in the present study, reflecting the typically conservative
 518 estimates adopted in building codes [38]. Similarly, NDMA [66], reported values ranging from

519 0.15g to 0.35g; the values obtained from the present study are found to lie near the lower to mid-
520 range of this bound. However, Pallav *et al.* [51] reported values between 0.13g and 0.19g, notably
521 lower than the findings of the present study. Overall, the present study's estimate is consistent with
522 prior research, reflecting a moderate seismic hazard assessment that is slightly higher than some
523 conservative estimates but lower than broader regional assessments.

524

525 **8. Summary and Conclusions**

526 This study explores the seismic hazard of Itanagar city, located in the state of Arunachal Pradesh
527 in Northeast India, using a probabilistic analysis approach. The analysis considers 18 fault lines
528 that contribute to ground motion within and around the study region. The output consists of PHA
529 and spectral acceleration (S_a) values for different return periods, aiding in seismic hazard
530 assessment and infrastructure design for the region. This study uses an earthquake catalogue from
531 1900-2024, consisting of 2054 mainshock events, to understand earthquake phenomena in the
532 region. Declustering is used to remove foreshocks and aftershocks, and the completeness of the
533 catalogue is examined using the CUVI method. Characterization of earthquake sources involves
534 developing tectonic maps, identifying all sources that can cause damaging earthquakes, measuring
535 fault lengths, determining the magnitude of the most damaging earthquakes observed, and
536 estimating the magnitude potential of seismogenic sources. The Geological Survey of India
537 developed SEISAT to analyze seismic activity in India. A 500 km radius around Itanagar is
538 examined using high-resolution scans of sheets 13-17, and a tectonic map is created. 33 active
539 tectonic features are identified, and 18 major active tectonic features capable of producing
540 significant ground motion were selected for seismic hazard analysis. Based on the present study,
541 the following conclusions are drawn:

- 542 • The b -value of Itanagar city ranges between 0.68 ± 0.04 to 0.89 ± 0.03 . These values reflect
543 the stress regime and tectonic complexity of the region.
- 544 • The GMPE proposed by Boore *et al.* [86] is identified as the most suitable model for the
545 region, enabling precise calculation of seismic hazard parameters. PHA values were
546 determined as 0.22g for a 2% probability of exceedance and 0.36g for a 10% probability
547 of exceedance, both within a 50-year timeframe.
- 548 • The spectral accelerations were computed for two return periods 475 and 2475 years at
549 specific time periods of 0.1 s, 0.3 s, 0.5 s, 1.0 s, 2.0 s, and 3.0 s. For the 475 year return

550 period, the spectral acceleration values are 0.54g, 0.38g, 0.25g, 0.125g, 0.048g, and 0.024
551 g, respectively. For the 2475 year return period, the corresponding spectral accelerations
552 are 0.95g, 0.80g, 0.48g, 0.250g, 0.09g, and 0.054g, respectively.

553
554 The results of the seismic hazard study will provide the necessary data to accurately identify and
555 map the seismically active zones, assess the seismic hazard, and determine the seismic intensity
556 levels in the area. This data can then be used to make informed decisions about seismic-resistant
557 design of structures, as well as seismic zonation studies, to ensure that structures are adequately
558 protected from seismic hazards. Further, future studies can be undertaken to improve magnitude
559 conversion methods by developing region-specific magnitude conversion equations for the study
560 area. Moreover, attempts should be made to detailed geological data, rupture lengths, and fault
561 displacement characteristics to produce more accurate seismic hazard assessments. Additionally,
562 the development of region-specific Ground Motion Prediction Equations (GMPEs) to achieve
563 precise ground-motion predictions and enable comparisons with global models for the Itanagar
564 region should be one of the prime focus of future studies.

565
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569
570 **Compliance with Ethical Standards**

571 **Conflict of Interest:** The authors declare that they have no known competing financial interests
572 or personal relationships that could have appeared to influence the work reported in this paper.

573 **Ethical Approval:** This article does not contain any studies with human participants or animals
574 performed by any of the authors.

575 **Informed Consent:** For this type of study, formal consent is not required.

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577 **JT:** Supervision, Formal analysis, Revision and Editing of drafted manuscript; **SSK & AD:**
578 Conceptualization, Revision and Editing of drafted manuscript

579

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